

Paleoclimatic Significance of Grain-size from Lacustrine Sediments in China

Huayong Li, Qingzhong Ming[#], Hucai Zhang, Lizeng Duan, Ziqiang Zhang

Laboratory of Plateau Lake Ecology and Global Change, Key Laboratory of Plateau Geographic Process and Environment Changes, College of Tourism & Geography Science, Yunnan Normal University, Kunming Chenggong 650500, China

[#]Email: mingqingzhong01@163.com

Abstract

In the research of using lacustrine sediments to recovery regional paleoclimate and paleoenvironment, the indicative significance of grain-size often has multiple interpretations. By summarizing the research results of 50 lacustrine sediments in China and combining the grain-size data of Gahai Lake and Langbapu Lake, we discuss the reason that the grain-size has multiple interpretations and the regular pattern of grain-size in different kinds of lakes. No matter in the short-time-scale and high-resolution research or in the long-time-scale and low-resolution research, the interpretations of grain-size for paleoclimate and paleoenvironment all have multiplicities. However the coarse sediments in the former indicates humid climate, the interpretation of the later one needs to consider many influence factors such as lake types and characteristics, the position of sampling points, the source of sediments and so on. If the sampling point is located in the deep-water area of the lakes, which are big and deep enough, the coarse sediment is probably dominated by high precipitation or strong wind. The three rings model of Grain size can affect the central area in the small and shallow lakes, meanwhile, the coarse grain content can also represent the strong precipitation, therefore, and the indicative significance of grain-size often has multiple interpretations in these kinds of lakes. Generally, coarse grain which lasts for a long period can represent the dry climate, which increase abruptly indicates high precipitation. Sampling points also have obvious influence on grain-size. Hydrodynamic force in the central of lakes and sediments source is simple relatively. Sediments are probely affected by runoff of entering into lake, aeolian and other factors. So it is more complex to explain.

Keywords: Grain size; Lakes; Paleoclimate; Paleoenvironment; Langpapu Lake; Gahai Lake

中国湖泊沉积物中的粒度变化及其古气候和古环境指示意义*

李华勇, 明庆忠, 张虎才, 段立曾, 张自强

云南师范大学 旅游与地理科学学院高原湖泊生态与全球变化实验室, 高原地理过程与环境云南省重点实验室, 云南 昆明 650500

摘要: 利用湖泊沉积物恢复区域内古气候和古环境的研究中, 粒度的指示意义往往具有多种不同的解释。本文通过总结我国近 50 个湖泊的粒度研究成果, 结合柴达木盆地尕斯库勒湖和元谋盆地浪巴铺古湖的研究数据, 探讨粒度具有多解性的原因, 以发现粒度在不同类型湖泊中的变化规律。无论是在短时间尺度、高分辨率, 还是在长时间尺度、低分辨率研究中, 粒度对于古气候和古环境的解释都具有多解性, 但前者粒度的变粗主要指示了降水增多的湿润气候, 后者粒度参数的解读要综合考虑湖泊类型与特点、采样点位置、沉积物来源等多重因素的影响; 面积和水深均较大的湖泊, 如采样点

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关键词: 粒度; 湖泊; 古气候; 古环境; 浪巴铺古湖; 尕斯库勒湖

粒度是利用湖泊沉积物重建古气候和古环境的常用代用指标之一, 因其具有样品需求量少, 测试简单、迅速、经济, 结果重复性好, 对气候变化敏感, 响应迅速等特点, 在古气候和古环境研究中得到广泛应用^[1-81]。前人研究发现, 受时间分辨率^[1, 2], 沉积物来源^[3, 4, 5, 6, 7, 8, 9, 10, 11, 12, 13, 14, 15, 16], 湖泊特征^[17, 18], 流域植被类型^[14, 19, 20], 采样点位置^[21, 22]等多重因素影响, 粒度参数对古气候和古环境的解释具有不确定性, 因而需借助其它指标对其指示意义进行约束。本文分析了我国近 50 个湖泊(古湖泊)^[1-81]的粒度研究结果, 结合本人的一些研究工作, 探讨不同条件下控制粒度参数(主要是中值粒径, 粘土和砂含量)发生变化的主导因素及其内在机制, 总结粒度在时空二维度中的变化规律, 力图解决粒度在解释古气候和古环境中存在的问题, 为今后的研究提供参考。

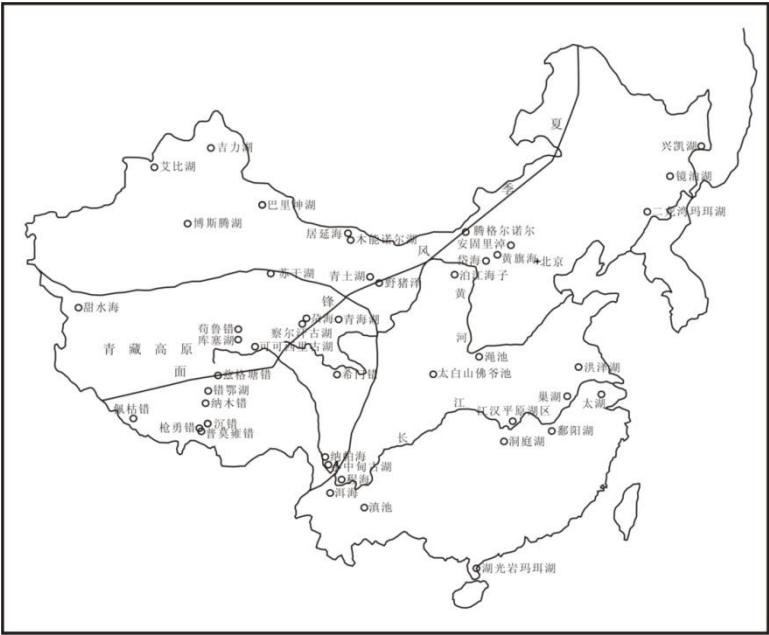


图 1 本文分析的中国主要湖泊分布图

1.1 短时间尺度

有学者研究洱海^[1, 23]、程海^[1]、滇池^[2]近现代以来湖泊沉积物中的高分辨率粒度变化, 对比气象观测数据发现, 沉积物粒径与流域降雨量变化具有较好的正相关关系, 两者的波峰和波谷能够分别一一对应, 说明粗颗粒组分的增加指示降雨量较大的湿润年份, 细颗粒组分的增加指示降雨量相对较小的干旱年份。因此得出结论, 在短时间尺度、高分辨率(年际、10a)研究中, 粗粒沉积物指示降雨量较大的湿润年份; 细粒沉积物指示降雨量相对较小的干旱年份。在最近几十年甚至是几百年内, 降水量的多寡不足以使湿润区湖泊发

生较大幅度的扩张和收缩^[1, 18, 23]，再加上有些湖泊本身具有出流口，限制了湖面的高度^[18, 23]。由此我们可以理解，在较短时间尺度内，相当一部分湖泊的湖面高度是比较稳定的，采样点相对于湖岸的距离没有发生太大变化，水动力条件也基本恒定。按照理想的湖泊沉积作用模式，从湖岸到湖心，随着水深的不断加大，其水动力条件由强变弱，湖泊沉积物具有从湖岸到湖心大致依次出现砾—一砂—粉砂—粘土的环带状分布规律^[24, 25]。在湖面保持基本稳定的时期内，粒度的同心圆式分布也基本不会发生扩展或收缩，因此流域内地表径流和地下渗流强度成为控制采样点粒度发生变化的最主要因素^[1, 2, 17, 18, 23]。湖泊汇水区降雨量大时，地表径流发育，较强的机械搬运作用可将较粗的陆源颗粒输入湖盆并沉降，从而导致沉积物颗粒变粗；降水量小时湖泊流域内地表径流不发育，携带能力减弱，只有少量陆源细颗粒能够到达采样点，沉积物粒径较小^[1, 2, 17, 23]。

即便在短时间尺度和高分辨率下，有学者研究安固里淖^[26]400年来和青藏高原苟鲁错^[27]最近半个世纪以来的沉积物粒度变化后发现，粗粒物质增多对应低湖面时的干旱期，与洱海^[1]、滇池^[2]、程海^[1]的研究结果正好相反。这种差异可能是由湖泊所处气候区以及自身特点的不同导致的。

1.2 长时间尺度

长时间尺度中湖泊沉积物粒度对古气候的反馈具有多种表现形式^[21, 28, 29, 30, 31, 32]，有些甚至截然相反^[33, 34]，这是由于在较长的时间跨度中，影响粒度变化的湖面高度^[35]，温度^[15, 36]，降水^[37]，风尘^[3, 4, 5, 6, 7, 8, 9, 10, 11, 12, 13]，植被^[38]，冰川^[15, 16, 39]等因素都可能发生较大的变化，而且研究者选取的钻孔位置，也会对其产生影响^[21, 40]。目前多数学者研究认为，气候干旱期，湖水位下降，采样点更靠近湖岸，水深变浅，水动力增强，因而沉积物颗粒较粗。相反，在气候湿润期，湖泊面积扩大，采样点距离岸边的距离较远，粗颗粒物难以到达，因而沉积颗粒较细^[26]。即：细粒沉积物指示降水较多的湿润气候，粗粒沉积物标志降水较少的干旱气候^[18, 25, 26, 27, 28, 29, 34, 35, 36, 38, 39, 40, 41, 42, 43, 44, 45, 46, 47, 48, 49, 50, 51, 52, 53, 54, 55, 56, 57, 58, 59, 60, 61, 62, 63]。

但在新疆吉力湖^[64]、博斯腾湖^[30, 37, 65]，东北二龙湾玛珥湖^[14]、镜泊湖^[17]，内蒙古木伦诺尔湖^[66]，青藏高原普莫雍错^[67]，云南点苍山冰川湖^[33]的研究中发现，粗粒沉积物增多时，反映流域内降水增多，地表径流水动力增强，有更多相对较粗的颗粒从流域地表被带到采样点沉降下来。当沉积物中细颗粒组分增多时，反映湖泊流域内气候干旱，降水稀少，地表水动力减弱，带入湖泊的颗粒以细颗粒为主^[14, 17, 23, 30, 31, 33, 37, 64, 65, 66, 67, 68]。

青海湖^[21]和兴凯湖^[32]的研究结果表明，粒度对古气候和古环境可能存在一种更为复杂的反馈关系。冷干和暖湿气候条件下，青海湖QH-2000钻孔中粗颗粒组分均表现为明显的增多。冷干气候条件下植被的减少，湖区水土保持能力的减弱，以及相对湖水水位的降低是造成入湖粗颗粒物明显增多的主要原因；暖湿气候条件下冰融水的大量补给或降水量的增加，导致湖区地表径流的相对发育，是造成入湖粗颗粒物明显增多的主要原因；介于二者之间的气候条件下，沉积物的粒径变化相对平缓。兴凯湖XK-1钻孔沉积物粒度揭示，粗粉砂和砂增多对应降水减少的低湖面时期，细粉砂增多对应于降水增加的高湖面时期，粘土增加对应气候干燥的静水沉积环境。

长时间尺度的湖泊沉积物中，粗粒沉积物的增多指示湖面收缩的干旱时期还是地表径流加强的湿润时期，需要结合不同对象的具体情况来具体分析，同时还需要参考其它古气候指标。一方面，当气候湿润，湖面升高，同一沉积位置由于距湖岸变远，湖水能量降低使得物质较细^[56]。另一方面，由于气候湿润条件下，降雨量大，地表径流增大，侵蚀能力增强，湖泊流体水动力增大，较粗颗粒也可以被带到离湖岸更远的地方，甚至湖心^[67]；反之，气候干旱时，可以是因湖面降低而沉积粗颗粒^[29]，而较弱的地表径流搬运粗颗粒物质的能力也会下降^[37]。通过分析前人研究成果，结合本人的一些工作，我们总结，粒度对古气候反馈形式的不同，主要与湖泊类型^[17, 28, 35, 37, 62, 67]、湖泊面积与水深^[30, 32, 39, 49, 67]、采样点位置^[21, 40, 67]等因素有关。

2 不同类型湖泊中的沉积物粒度记录

2.1 封闭与开口湖泊

无论是粗粒还是细粒沉积物的增多指示高湖面时期，其前提条件之一是湖盆必须在一个相对较大的空间内是封闭的，这样才可能出现历史时期的高湖面。但在开口湖中，粒度的记录则恰好相反。侯新花等^[17]研究了镜泊湖（开口湖）全新世以来的粒度变化后发现，湖泊沉积物主要由入湖河流和降水冲刷携带而来，在水动力较强时，细颗粒来不及沉积下来就随着出湖水流动走了，只有粗颗粒沉积下来，而在水动力较弱时，大部分细粒颗粒物质沉淀在湖中。因此开口湖中粗颗粒的增多反映流域降雨量较大，湖水水动力较强，细粒物质的增多对应降雨量较小的干旱时期，与封闭湖泊中的记录恰好相反^[63]。

2.2 不同面积和水深

湖泊面积和水深决定了其面对气候变化时的稳定性，进而决定了湖心沉积物粒度受湖泊扩大和收缩影响的大小。同处我国内陆干旱区的博斯腾湖^[65]和巴里坤湖^[41]，相同的粒度特征却指示完全相反的古气候状况，可能是由两者不同的特征决定的。博斯腾湖湖水面积约 1000 km²，平均水深 8 m，最大水深 16.2 m，采样点位于湖心最大水深处，至今没有发现高湖面的湖岸堤或湖岸阶地，表明湖面波动不大，湖泊中心沉积物的粗细主要取决于入湖水动力和流量的大小。巴里坤湖湖面面积不足 100 km²（2005 年），平均水深 0.6 m，最大水深约 1 m，古湖岸线显示巴里坤湖面积曾达到约 800 km²，1984 年湖水面积萎缩至 112 km²，研究表明粗颗粒指示了较低的湖水位，气候较为干旱，物理风化强盛，植被覆盖度下降；细粒沉积物增多指示了较高的湖水位，气候较为湿润，地表化学风化相对增强，植被盖度增加。

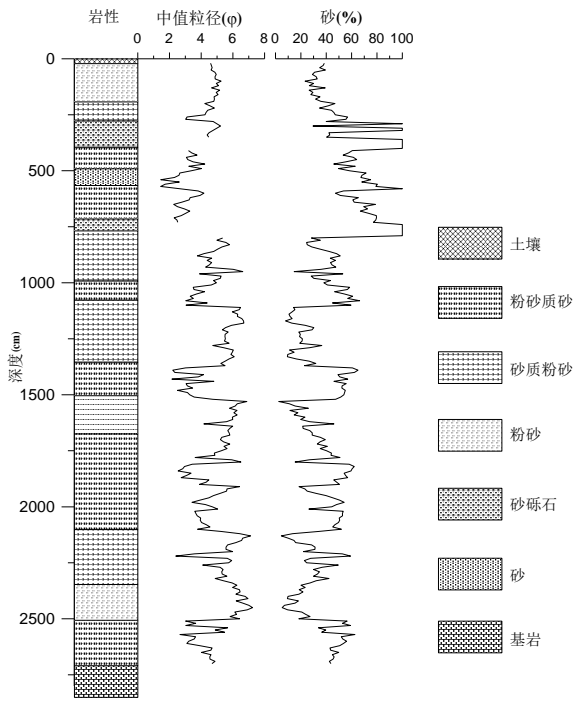


图 2 浪巴铺剖面粒度曲线

有学者提出，对于面积和水深均较大的湖泊，很难出现湖心水深浅到能够受到表层水流扰动的条件^[67]。因此对于这类湖泊，其湖心沉积物粒度的变粗主要因为流域降水增加，地表侵蚀能力加强，相对较粗的颗粒被带入湖盆^[31, 65, 66, 67, 68]。而对于面积和水深均较小的湖泊^[41, 55, 62]，湖面容易发生明显的扩张和收缩，湖心离岸较近时，沉积物颗粒变粗，湖心离岸较远时，沉积物颗粒变细。

2.3 古湖泊——元谋盆地浪巴铺古湖粒度特征及气候意义

浪巴铺古湖位于云南元谋县城以西 20km 的浪巴铺村附近，地理坐标 25°41'5" N, 101°42'2" E, 海拔 1630m。古湖面积约 1km²，从周围地貌来看，湖泊深度和汇水面积都十分有限，因此湖泊水位易受降水影响，发生较

大幅度变化。2.8~4.0 m、5.5~6.0 m、7.3~8.1 m 等 3 处为砂砾石层，分布连续，磨圆度高，最大粒径 2 cm，为典型的湖滨相沉积。由此可判断浪巴铺古湖曾收缩到剖面位置，因此浪巴铺古湖沉积物中细颗粒沉积物增加指示湖面上升的湿润时期，粗颗粒沉积物增加指示水位下降的干旱时期。浪巴铺剖面沉积物中值粒径(ϕ)变化范围为 1.45~7.24，中值粒径为 4.67。砂含量 2.58%~100.00%，平均含量 42.86%。黏土含量 0.00%~17.80%，平均含量 7.82%。从图 2 可发现，浪巴铺古湖经历了至少 7 次较大幅度的湖面涨落^[88]。

3 同一湖泊的不同采样点位置对粒度测试结果的影响

同一湖泊的不同部位，其粒度变化规律可能截然不同，这在一些面积较大的湖泊中表现尤其明显^[21, 40, 69]。总结前人研究不难发现，在湖岸阶地^[56, 62]、古湖相沉积物^[29, 46, 50, 54]、离岸较近^[55]、湖滨陆地^[35, 41, 42, 57]取样，得出的结论多是粗粒沉积物的增加指示低湖面时的干旱时期，细颗粒的增加对应高湖面时的湿润期。因为湖岸线附近，是湖面扩张和收缩影响最直接的区域，而且古湖岸阶地^[55, 56, 62]和古湖岸线^[41, 62]，也留下了确凿的湖面涨落证据。湿润期湖面升高，采样点远离湖岸，沉积物粒度变细，反之沉积物粒度变粗。

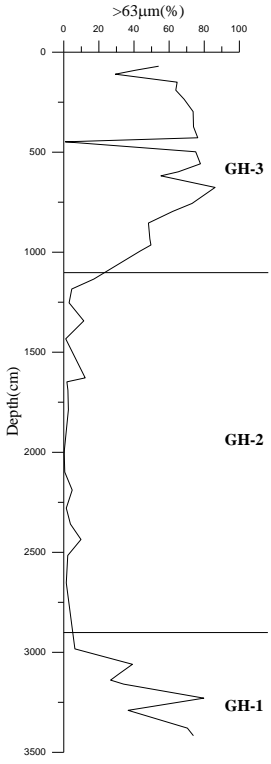


图 3 尕斯 GH-2 孔砂含量变化曲线

2003 年 5 月，中科院青海盐湖所用汽车钻在尕斯湖西岸钻取 DG02 和 DG03 岩芯，地理坐标为 DG02 钻孔 37°10' N, 97°33' E, DG03 钻孔 37°08' N, 97°33' E。DG02 钻岩芯长 34.57 m，在不同层段选取 48 个样品进行粒度分析，结果见图 3。岩芯按粒度特征分为三段，GH-1 阶段沉积物粒度组成以砂和粉砂为主，孢粉资料^[35]和相邻 DG03 钻孔的数据^[42, 82-87]表明，这段时期气候干旱但不稳定，钻孔点水位较低，沉积物粒径为整段岩芯中较粗的一段。GH-2 阶段粗粒组分含量非常低，主要以粘土和粉砂为主，为全剖面最细的一段。其他代用指标^[82-87]也表明，这段时期是全新世气候最为湿润的一段时期，气候进一步湿润，湖面上升，钻孔位置水深较大，距离湖岸位置较远，沉积物以细粒物质为主。GH-3 阶段沉积物组成以砂和粉砂为主，砂含量是全剖面最高的一段，孢粉分析显示该段代表干旱半干旱生境的藜科花粉含量显著增加^[35]，表明该段气候明显变冷，降水也大为衰减，湖泊强烈收缩，湖面下降，钻孔位置水深剧烈减小甚至出露在湖面以上，沉积物以粗粒物质为主。

钻孔地点选在湖心^[64, 65]或者靠近湖心的深水区^[14, 67], 如果其湖泊面积和水深足够大, 则湖面的扩张和收缩一般不能对中心部位的沉积物粒度产生直接影响, 因此湖中心粒度的变化主要受区域内降水量和降水强度的影响。殷志强等^[24]分析我国北方内陆湖泊安固里淖湖泊沉积物粒度特征, 将其划分为湖滨带, 湖滨-湖心过渡带, 湖心带, 见图 4。安固里淖是河北坝上地区最大的内陆封闭湖泊, 面积 47.6 km², 最大水深 4.0 m, 平均水深 2.5m, 湖盆呈浅碟状, 湖底平坦。根据殷志强等的研究结果和图 3 可知, 从湖岸到湖心, 粗颗粒物逐渐减少, 细粒物质增多, 湖心区所占面积最大, 粒度呈现以 40 μm 为峰值的单峰模式。粒度的这种环带状分布规律在中国多数湖泊中都是存在的^[25]。我们注意到, 湖滨的砂带和湖滨-湖心过渡的粉砂带, 宽度十分有限, 湖盆以湖心粘土带为主, 如果钻孔地点选在湖泊中心深水区, 那么较小规模的湖泊收缩和扩张, 不能对其粒度特征产生影响, 而一次较强的降雨^[14, 30, 31, 33]或者风沙过程^[71, 72], 则可以将相对较粗的颗粒带至湖心并沉积下来。此时沉积物粒度的变粗指示了降水较多的湿润气候或者沙尘较多的恶劣气候。这种结论在中国不少湖泊研究中得到了印证^[14, 64, 65, 66, 67, 68]。

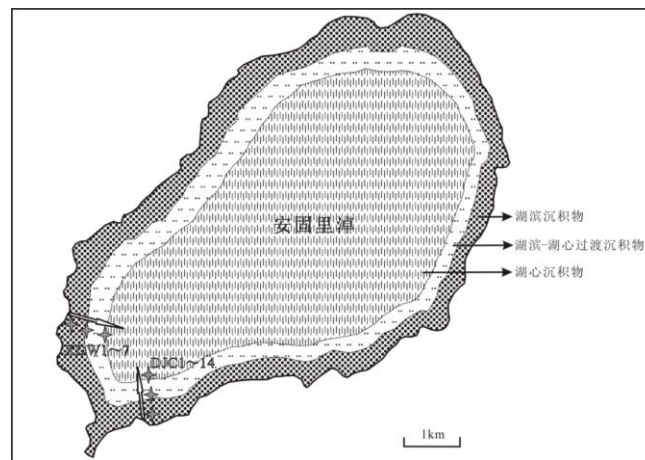


图 4 安固里淖湖泊沉积物环带状分布规律 (据殷志强等, 有修改)

4 结论

(1) 无论是在短时间尺度、高分辨率研究, 还是在长时间尺度、低分辨率研究中, 粒度对于古气候和古环境的解释都具有多解性。但前者粒度的变粗主要指示了降水增多的湿润气候, 后者粒度参数的解读要综合考虑湖泊类型与特点、采样点位置、沉积物来源等多重因素的影响。

(2) 对于面积和水深均较大的湖泊, 如采样点位于湖心深水区, 其沉积物粒径变大应首先考虑流域地表侵蚀增强或者“尘暴”爆发导致沉积物来源的变粗, 而非干旱气候下的湖泊萎缩事件。

(3) 面积小而且水深较浅的湖泊, 粒度的环带状分布模式影响可大至湖心, 但同时必须注意到, 降水增多时, 地表侵蚀加剧, 同样能够把更多粗颗粒带入湖泊甚至湖心, 因此这类湖泊粒度的解释存在不确定性。一般可理解为较长时间跨度的粒度变粗和变细分别对应湖泊收缩和扩张, 而粒度的突然变粗则指示短时间内的降水增多。

(4) 我国西北地区有很多已经消亡和正在逐步消亡的湖泊, 事实证明该地区古湖曾广泛发育, 并发生巨大的湖面涨落。该区域湖相和古湖相沉积物的变粗反映了湖面下降, 湖泊萎缩事件, 但需要甄别沉积物中水成和风成组分。

(5) 采样点的选择对粒度的影响也是显著的, 相比来说, 湖心水动力和沉积物来源单一, 解释也相应单一, 而湖滨沉积物则可能受到河流三角洲、风沙活动、浊流等多种因素影响, 解释起来更加复杂。

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【作者简介】



李华勇（1986-），男，山东寿光人，博士研究生，研究方向为湖泊沉积与气候变化。

Email: lihuayong2010@hotmail.com

明庆忠（1963-），男，湖北黄冈人，教授，博士生导师，主要从事区域地理环境与区域发展研究。

Email: mingqingzhong01@163.com